

Microbially mediated CH₄ consumption and N₂O emission is affected by elevated CO₂, soil water content, and composition of semi-arid grassland species

Feike A. Dijkstra · Jack A. Morgan ·
Daniel R. LeCain · Ronald F. Follett

Received: 27 February 2009 / Accepted: 24 August 2009 / Published online: 2 September 2009
© Springer Science + Business Media B.V. 2009

Abstract Elevated CO₂ affects plant productivity, but also water availability and plant species composition in semi-arid grasslands, thereby potentially causing complex effects on CH₄ consumption and N₂O emission. We studied the effects of atmospheric CO₂ concentration (400 vs 780 μL L⁻¹), water content (15 vs 20% gravimetric soil moisture), and composition of semi-arid grassland species (perennial grasses *Bouteloua gracilis*, *Hesperostipa comata*, and *Pascopyrum smithii*; sub-shrub *Artemisia frigida*; invasive forb *Linaria dalmatica* grown in monoculture and all five species together) on CH₄ consumption and N₂O emission in a full factorial greenhouse experiment. We used a unique method where we measured microbial effects on CH₄ consumption and N₂O emission in isolation from effects of gas diffusivity. Microbially mediated CH₄ consumption was significantly higher under elevated CO₂ (by 20%), but was not affected by soil water content or plant species

composition. Microbially mediated N₂O emission was not significantly affected by elevated CO₂, but was significantly higher with high water content (by 67%) and differed significantly among species. Treatment effects on CH₄ consumption and N₂O emission often could not be explained simply by differences in soil moisture, suggesting that treatment-induced changes in other soil and microbial properties played a role in causing these effects.

Keywords Elevated CO₂ · GRACenet publication · Methane · Microbial activity · Nitrous oxide · Soil moisture

Introduction

Earth's atmospheric CO₂ concentration has risen from ~280 μL L⁻¹ at the start of the industrial revolution to greater than 385 μL L⁻¹ today, and is expected to exceed 700 μL L⁻¹ by the end of this century (Intergovernmental Panel on Climate Change 2007). This rise in atmospheric CO₂ has strong direct and indirect effects on ecosystems. Semi-arid grasslands are considered to be highly responsive to rising CO₂. Biological activity in these grasslands is strongly modulated by the natural seasonal variation in soil water availability (Frank and Groffman 1998; Huxman et al. 2004; Potts et al. 2006). Elevated CO₂-induced increases in plant water use efficiency due to increased plant stomatal closure (Morgan et al. 2004b) would be

Responsible Editor: Elizabeth M. Baggs.

F. A. Dijkstra (✉) · J. A. Morgan · D. R. LeCain
USDA-ARS, Rangeland Resources Research Unit,
Crops Research Laboratory,
1701 Centre Ave,
Fort Collins, CO 80526, USA
e-mail: feike.dijkstra@ars.usda.gov

R. F. Follett
USDA-ARS, Soil, Plant, and Nutrient Research Unit,
2150 Centre Ave,
Fort Collins, CO 80526, USA

expected to enhance that biological activity through increases in soil water content. Indeed, CO₂-induced changes in soil water availability may have played an important role in the observed increases in plant productivity, shifts in plant species composition, and altered C and N cycling in semi-arid grasslands exposed to enriched CO₂ concentrations (Dijkstra et al. 2008; Morgan et al. 2004b, 2007; Pendall et al. 2003).

A major concern of a global increase in atmospheric CO₂ is how it will affect flux rates of other greenhouse gases such as methane (CH₄) and nitrous oxide (N₂O). Methane and N₂O have a 25 and 298 times greater relative global warming potential than CO₂ over 100 years (Intergovernmental Panel on Climate Change 2007). Semi-arid grasslands represent a significant global sink for CH₄ and source for N₂O (Galbally et al. 2008; Mosier et al. 1991, 1997). The long-term effects of elevated CO₂ on CH₄ consumption and N₂O emission in these grasslands are largely unknown (Galbally et al. 2008), in part because of complicated effects of CO₂ on soil water, C and N availability, as well as plant productivity and species composition.

Changes in atmospheric CO₂ concentration can affect CH₄ consumption and N₂O emission in semi-arid grasslands through changes in soil water availability. Methane consumption often decreases in response to elevated CO₂, which has been related to increased soil moisture (Ambus and Robertson 1999; Ineson et al. 1998; McLain and Ahmann 2008; McLain et al. 2002), but no change in CH₄ consumption has also been observed (Kang et al. 2001; Kettunen et al. 2005; Mosier et al. 2002). Although a certain amount of soil water is required for methanotrophs to oxidize CH₄, influxes of CH₄ into soils may be impeded once soil water content exceeds threshold values beyond which diffusivity of CH₄ becomes more limiting (Del Grosso et al. 2000; Koschorreck and Conrad 1993). On the other hand, increases in soil water content have been implicated in CO₂-related increases in N₂O emission (Arnone and Bohlen 1998; Kanerva et al. 2007; Kettunen et al. 2006; Robinson and Conroy 1999). Elevated CO₂ may also elicit changes in CH₄ and N₂O exchange unrelated to soil moisture. Elevated CO₂ could decrease CH₄ consumption by transforming the microbial communities that are responsible for CH₄ production and consumption (McLain and Ahmann

2008; Phillips et al. 2001a), while this decrease in CH₄ consumption may depend on elevated CO₂-induced changes in gross nitrification (Baggs and Blum 2004). Others have suggested that greater N₂O emissions under elevated CO₂ are a consequence of enhanced rhizodeposition, acting as an energy source for denitrification (Baggs et al. 2003a, b; Ineson et al. 1998; Kammann et al. 2008). But often N₂O emissions do not respond to elevated CO₂ (Ambus and Robertson 1999; Billings et al. 2002; Hungate et al. 1997a; Mosier et al. 2002). Mosier et al. (2002) suggested that a combination of greater N supply (due to increased N mineralization with increased soil moisture) and increased plant N demand under elevated CO₂ maintained mineral N pools in the soil too low to alter rates of N₂O emissions in a semi-arid grassland.

Elevated CO₂ could also potentially affect CH₄ consumption and N₂O emission in semi-arid grasslands through changes in plant species composition. Elevated CO₂ caused greater abundances of the C₃ grass *Hesperostipa comata* and the sub-shrub *Artemisia frigida* in a semi-arid grassland in Colorado (Morgan et al. 2004a, 2007). However, little is known about how such changes in species composition will affect CH₄ and N₂O fluxes. Epstein et al. (1998) reported greater CH₄ consumption in patches dominated by C₄ grasses (mostly *Bouteloua gracilis*) than in patches dominated by C₃ grasses (mostly *Pascopyrum smithii*) in a semi-arid grassland. Further, patches mixed with C₃ and C₄ grasses had lower N₂O emissions than either the C₄ or C₃ dominated patches. Niklaus et al. (2006) also reported lower N₂O emissions with increased diversity of temperate grassland species. In a sagebrush-steppe Norton et al. (2008) reported greater N₂O emissions under the exotic grass *Bromus tectorum* than under the native grass *P. smithii*. Plant species effects on CH₄ consumption and N₂O emission have been associated with plant species effects on microbial community composition (Menyailo and Hungate 2003; Ullah et al. 2008), N cycling and soil NO₃⁻ concentration (Crenshaw et al. 2008; Epstein et al. 1998; Norton et al. 2008), and on soil C:N ratios (Ambus et al. 2006; Menyailo and Huwe 1999).

Responses of CH₄ consumption and N₂O emission are often complex because microbial and physical factors (gas diffusivity) controlling these fluxes operate at different scales in time and space.

We set up a controlled greenhouse experiment to better understand microbially mediated CH₄ consumption and N₂O emission. In particular, we studied how elevated CO₂, soil water content, species composition of semi-arid grasslands, and their interactions affected microbially mediated CH₄ consumption and N₂O emission in isolation from physical processes (i.e., gas diffusivity). In two greenhouses we grew five grassland species in monoculture and all five species in competition under ambient and elevated CO₂ (400 vs. 780 μL L⁻¹), and under ‘low’ and ‘high’ soil moisture conditions (15 vs 20% gravimetric soil moisture). We used the native perennial grasses *Bouteloua gracilis*, *Hesperostipa comata*, and *Pascopyrum smithii*, the sub-shrub *Artemisia frigida*, and the invasive forb *Linaria dalmatica*, all common to the semi-arid grasslands of Colorado. We used a novel method to measure CH₄ consumption and N₂O emission in which we eliminated potential effects of gas diffusivity. In this way we were able to test the direct effects of the treatments on microbially mediated CH₄ consumption and N₂O production. We addressed the following questions: 1) how do elevated CO₂, soil water content, and species composition affect microbially mediated CH₄ consumption and N₂O emission, and 2) to what extent can these effects be explained by soil moisture? Besides soil moisture, we also tested if parameters of plant activity (pot respiration, plant biomass and its N content) or inorganic soil N could explain the variation in CH₄ consumption and N₂O emission among treatments.

Materials and methods

Experimental design

The soil used in this experiment came from the USDA-ARS Central Plains Experimental Range in the shortgrass steppe region of north-eastern Colorado. The soil was a sandy loam of the Ascalon series (Aridic Argiustolls). The soil collected to 20-cm depth was carbonate-free, had a pH of 6.6 and 0.95% total C and 0.09% total N. The soil was homogenized by sieving (4 mm) and air-dried before use. We filled each of 96 polyvinyl chloride (PVC) pots (diam. 20 cm, height

40 cm, closed at the bottom except for an air inlet) with 14 kg of air-dried soil (Fig. 1). Before filling the pots we placed a nylon bag filled with 3 kg playground sand at the bottom of each pot to improve air circulation. Exact weights of soil and filled pots were recorded. We watered the pots to field capacity (i.e., 30% gravimetric soil moisture content). In each pot we put five plants of *Artemisia frigida* (sub-shrub), *Linaria dalmatica* (forb), *Bouteloua gracilis* (C₄ grass), *Hesperostipa comata* (C₃ grass), or *Pascopyrum smithii* (C₃ grass), 16 pots for each species. All species are native to the shortgrass steppe, except for the invasive weed *L. dalmatica*. We also planted 16 pots with all five species combined (one plant of each species per pot). Plants were raised from seeds in peat pellets (2–5 seeds in each peat pellet). After plant emergence, plants were thinned to one plant per peat pellet and transferred to the pots.

We grew the plants in two greenhouses at the USDA-ARS Crops Research Laboratory in Fort Collins, Colorado. In one greenhouse we raised the atmospheric CO₂ concentration to a constant level of 780±50 μL L⁻¹ (average ± standard deviation) by adding pure CO₂, while in the other greenhouse the atmospheric CO₂ concentration was kept near ambient level (~400±40 μL L⁻¹, note that this concentration is slightly higher than the average global CO₂ concentration because our greenhouses were in an urban setting). The CO₂ concentration was continuously monitored and the CO₂ supply was computer-controlled (Argus Control Systems Ltd, White Rock, BC¹). The added CO₂ entered the greenhouse through a ventilation system ensuring uniform distribution of the CO₂ concentration inside the greenhouse. Air temperature in both greenhouses was kept between 27 and 29°C during the day and between 16 and 18°C during the night. Temperature was regulated by computer-controlled air conditioners and heaters (York International, York, PA). Both greenhouses were equipped with 600 W lights (P.L. Light Systems, Beamsville, ON) that were on during the day for 12 hrs. The light intensity in each greenhouse was ~200 Wm⁻² during the day. The relative humidity in each greenhouse was 24±5% during the day and 30±5% during the night.

¹ Trade and company names are given for the reader's benefit and do not imply endorsement or preferential treatment of any product by the USDA.

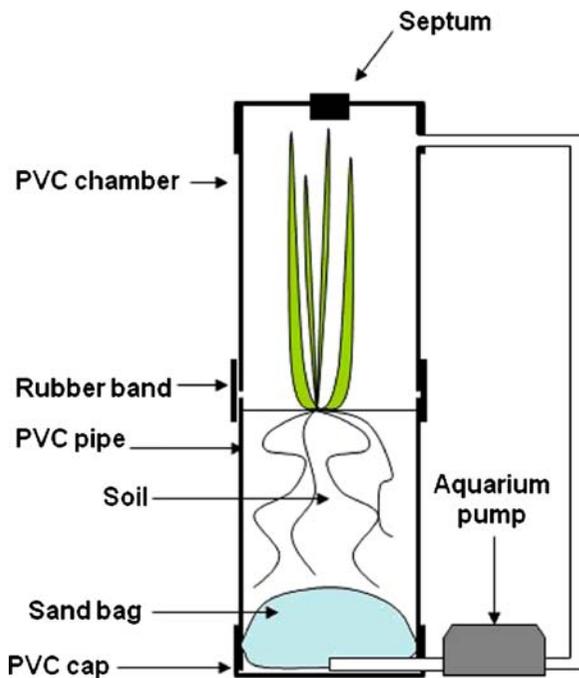


Fig. 1 Setup of the pots during CH_4 and N_2O flux measurements

Half of the pots (8 pots for each species and 8 pots with all species combined) were placed in the greenhouse under elevated CO_2 (total of 48 pots) and the other half of the pots were placed in the greenhouse under ambient CO_2 (total of 48 pots). Once every week, the pots were swapped between the two greenhouses and the CO_2 concentration switched accordingly. In this way we tried to minimize greenhouse effects and pseudoreplication at the CO_2 level (Goverde and Erhardt 2003; Heijmans et al. 2002).

The pots were watered frequently (see below) to restore gravimetric soil moisture content to 15% (low water) to half of the pots, and to 20% (high water) to the other half of the pots (4 pots/replicates for each CO_2 and species composition treatment), or 50 and 67% of field capacity respectively. During the first week, all pots were maintained at 30% soil moisture content, after which pots were allowed to dry down to their soil moisture target contents (15 and 20%). Once target soil moisture contents were reached (26 days after transplanting), pots were watered three times a week. Pots were weighed once every week and watered up to their target weights. The amount of water added during the other two times of the week was calculated based on previous water loss from each pot. In each greenhouse, pots were placed in four blocks, where one replicate of

each of the species composition and water treatments were randomly placed in each block (each block consisting of 12 pots).

Measurements and analyses

We measured CH_4 consumption and N_2O emission 48, 69, and 83 days after transplanting, using a modified method developed for respiration in pots in the presence of plants (Cheng 1996). We also measured respiration at the same time, and used respiration as a covariate to explain variation in CH_4 consumption and N_2O emission among treatments (see below). On each measurement day we placed chambers (diam. 20 cm, height 45 cm) fitted with a septum and air outlet on top of the pots (Fig. 1). We circulated air inside the pot/chamber by inserting an aquarium pump (Apollo AM-3, Apollo Enterprises, Ventura, CA, flow rate 2.8 L min^{-1}) in-line between the air inlet at the bottom of the pot and the air outlet at the top of the chamber. All pots/chambers were air-circulated for two hours to completely mix the air inside the chamber and pot. During this time CO_2 inside the pot/chamber was removed by an in-line CO_2 scrubber (PVC tube, diam. 3.5 cm, height 36 cm, filled with sodalime). We then removed the scrubber and pulled a 30 ml gas sample from each chamber (time 0) and another 30 ml gas sample two hours later while air still circulated inside the pot/chamber. Chambers and pumps were removed after the second gas sample was taken. Gas samples were analyzed for CH_4 , N_2O , and CO_2 on a gas chromatograph (Varian 3800, Palo Alto, CA). All species remained in a vegetative stage until the end of the experiment, except for *B. gracilis* and *L. dalmatica* that started to flower during the last measurement. Under field conditions, flowering of the perennials *B. gracilis* and *L. dalmatica* occurs throughout the growing season that does not stop their growth (indeterminate growth). Also in our experiment, *B. gracilis* and *L. dalmatica* kept growing after flowering.

The purpose of circulating air through the pot/chamber was to establish well-mixed and uniform concentrations of CH_4 and N_2O inside the pot/chamber thereby eliminating effects of gas diffusivity on CH_4 consumption and N_2O emission. Thus, the rate of CH_4 consumption and N_2O emission that we measured indicates the activity of the microbial community. The CH_4 consumption ($\mu\text{g C pot}^{-1}\text{hr}^{-1}$), N_2O emission

($\mu\text{g N pot}^{-1}\text{hr}^{-1}$), and respiration ($\text{mg C pot}^{-1}\text{hr}^{-1}$) were calculated based on the difference between the concentrations of the two samples and the air volume of the pot/chamber system. We assumed that the decrease in CH_4 concentration and increase in N_2O and CO_2 concentration with time did not affect their respective consumption and production rates. For N_2O and CO_2 production, this is a reasonable assumption, since N_2O and CO_2 concentrations in the soil tend to be much higher than in the atmosphere. However, for CH_4 consumption, this may not be true. The decrease in CH_4 concentration inside the pot/chamber may have reduced the rate of CH_4 consumption. However, the average CH_4 concentrations at 2 and 4 hrs after chambers were placed on the pots were 1.52 and 1.09 $\mu\text{l L}^{-1}$ respectively. When we assume a CH_4 concentration at the time when chambers were placed on the pots of $\sim 1.9 \mu\text{l L}^{-1}$ (i.e., ambient CH_4 concentration measured in the greenhouse after termination of the experiment), then our assumption of a linear decrease in CH_4 concentration with time inside the pot/chamber is reasonable. We also calculated a first order rate constant k to describe the rate of CH_4 consumption ($-kt = \ln(\text{CH}_{4,t1}/\text{CH}_{4,t0})$, Hütsch et al. 1994), and treatment effects on k were very similar to treatment effects on CH_4 consumption calculated with the linear approach (data not shown). We feel confident that treatment effects on our gas flux would not change if we added more data points in time. It would be more accurate to refer to *net* CH_4 consumption and *net* N_2O production, because of simultaneous CH_4 production (by methanogens) and consumption (by methanotrophs) and N_2O production (through nitrification and denitrification) and consumption (complete denitrification). However, because of the relatively low soil moisture content and high oxygen supply during our measurements we assumed CH_4 production and complete denitrification to N_2 to be negligible. We did not attempt to characterize the different microbial groups responsible for CH_4 and N_2O consumption and production. We are aware that our artificial way of measuring CH_4 and N_2O fluxes in a greenhouse pot study does not allow us to extrapolate our results quantitatively to field situations. However, we want to emphasize that the purpose of this study was to better understand microbially mediated CH_4 consumption and N_2O emission that are not confounded by gas diffusivity effects.

We harvested all plants after the final CH_4 consumption and N_2O emission measurement (85 days after

transplanting). Plants were separated into aboveground biomass, crowns, and roots. Plant biomass was dried (60°C) and weighed. Plant biomass was analyzed for N concentration on a mass spectrometer (20–20 Stable Isotope Analyzer, Europa Scientific, Chesire, UK). The soil from each pot was thoroughly homogenized before taking a sub-sample for analysis. We added 60 ml of 2 M KCl to 25 g moist soil. Samples were shaken for 1 h and filtered through pre-leached (with 2 M KCl) Whatman No. 1 filter paper. Extracts were frozen until analyses for NH_4^+ and NO_3^- on a flow injection analyzer (QuickChem FIA + , Lachat Instruments, Milwaukee, WI). The NH_4^+ and NO_3^- concentrations were expressed on oven-dry soil weight (105°C) basis.

We used repeated measures ANOVA to test for main effects of CO_2 (two levels: ambient and elevated), water (two levels: low and high water), species composition (six levels: five species grown in monoculture and all five species grown in competition), and their interactions on CH_4 consumption and N_2O emission. We also used repeated measures ANOVA to test for main effects of CO_2 , water, and species number (two levels: 1 and 5 species), and their interactions on CH_4 consumption and N_2O emission. The repeated measures ANOVA included random effects of date (48, 69, and 83 days after transplanting) and block. We also used ANOVA to test for main effects of CO_2 , water, species composition, and their interactions (or for main effects of CO_2 , water, species number, and their interactions) on CH_4 consumption and N_2O emission at each date. We then tested for the same effects after adjusting for effects caused by soil moisture (covariate) using ANCOVA. Treatment effects on CH_4 consumption and N_2O emission could be mediated by treatment effects on plant growth, inorganic N uptake, soil NH_4^+ and NO_3^- , and rhizodeposition. We therefore also used plant biomass, plant biomass N content (an integrated measure of plant N uptake), soil NH_4^+ and NO_3^- measured at the end of the experiment, and respiration (which includes rhizodeposition) as a covariate. We realize that soil NH_4^+ and NO_3^- measured at the end of the experiment may not reflect dynamics of their availability during the experiment. Also, respiration may be a poor indicator of rhizodeposition. Unfortunately, we have no direct measurements of soil NH_4^+ and NO_3^- and rhizodeposition during the experiment. Plant biomass, plant biomass N content, and soil NH_4^+ and NO_3^- were only used as a covariate for the analyses of the

last date of CH_4 and N_2O measurements. We used linear regressions to relate CH_4 consumption and N_2O emission to soil moisture, respiration, plant biomass, plant biomass N content, or NH_4^+ and NO_3^- . When necessary, data were log transformed to reduce heteroscedasticity. All statistical analyses were done with JMP (version 4.0.4; SAS Institute, Cary, North Carolina, USA).

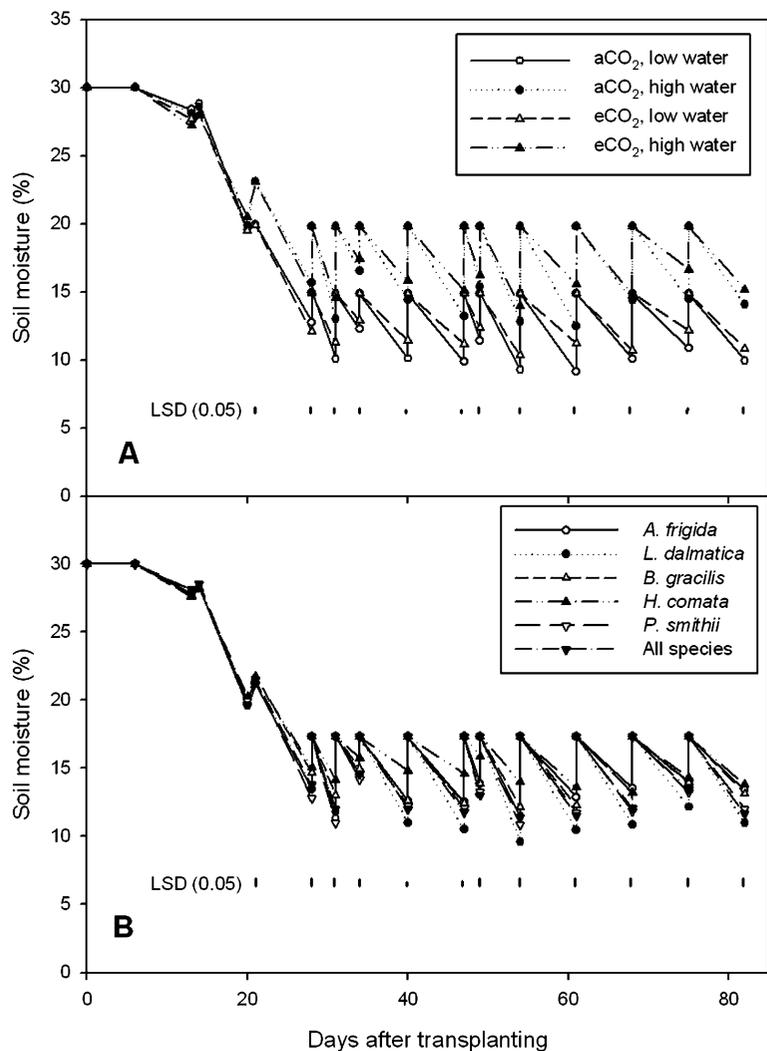
Results

Between watering events pots under ambient CO_2 dried out faster than pots under elevated CO_2 (Fig. 2a). Soil moisture contents also diverged among pots with

different species composition after each watering event (pots with *L. dalmatica* drying out fastest and pots with *H. comata* slowest, Fig. 2b). Because of differences in soil moisture among CO_2 and species composition treatments, we used soil moisture content as a covariate to test if treatment effects on CH_4 consumption and N_2O emission still existed after adjusting for soil moisture effects. We should note that the soil moisture content measured at the time of CH_4 and N_2O flux measurements may not reflect effects caused by differences in the drying-rewetting cycles that occurred throughout the experiment.

We observed CH_4 consumption in all pots at all three dates. Repeated measures ANOVA on all three dates showed that elevated CO_2 significantly in-

Fig. 2 Gravimetric soil moisture concentration during the experiment for **a** the CO_2 (a CO_2 : ambient CO_2 , e CO_2 : elevated CO_2) and water treatments (low water: 15% soil moisture, high water: 20% soil moisture), averaged across species composition treatment and **b** the species composition treatment averaged across the CO_2 and water treatments



creased the CH₄ consumption (on average by 20%), while there were no significant water and species composition effects (Table 1, Fig. 3). There were also no significant treatment interactions on CH₄ consumption. The increase in CH₄ consumption under elevated CO₂ was not significant during the first date when plants were still small, but became significant during the last two dates of measurement (ANOVA, Table 1, Fig. 3). After adjusting for soil moisture effects, CH₄ consumption was still higher under elevated CO₂ during the last two dates of measurement (ANCOVA, Table 1). Surprisingly, we observed no significant relationship between the CH₄ consumption and soil moisture content for each date or when all three dates were combined ($P>0.1$, Fig. 4a). Methane consumption 83 days after transplanting was not significantly related to plant biomass, plant biomass N, or NH₄⁺ and NO₃⁻ concentrations in the soil ($P>0.1$) and adjusting for plant biomass, plant biomass N, or NH₄⁺ and NO₃⁻ concentrations in the soil did not change treatment effects (data not shown). Methane consumption was significantly positively related to pot respiration, although the variability explained was small ($P<$

0.0001, $R^2=0.18$). When respiration was used as a covariate in the ANCOVA it did not change treatment effects (data not shown).

The emission of N₂O at all three dates was not significantly affected by elevated CO₂, but was significantly higher in the high water treatment (on average by 67%) and differed significantly among the six levels of the species composition treatment (repeated measures ANOVA, Table 1, Fig. 5). The N₂O emission was highest under *H. comata*, particularly in the high water treatment (significant water*species composition interaction). The ANOVA results for each date show that water treatment effects became less significant by the end of the experiment, but that species composition effects remained significant throughout the experiment (mostly due to *H. comata*). When we used soil moisture content as a covariate, the water treatment effect disappeared 48 and 83 days after transplanting and was only marginally significant 69 days after transplanting. However, species composition effects remained significant for all dates and the water*species composition interaction remained the same after correcting for soil moisture effects (ANCOVA, Table 1). Not surprisingly, the N₂O

Table 1 Summary of statistic test results with CO₂, Water, and Species composition as main factors^a

Test	Covariate	P-values							
		CO ₂	Water	Sp. comp.	CO ₂ * Water	CO ₂ * Sp. comp.	Water* Sp. comp.	Covariate	
CH ₄									
Rep. meas.	ANOVA	—	<0.0001	ns	ns	ns	ns	ns	—
48 DAT	ANOVA	—	ns	ns	ns	ns	ns	ns	—
69 DAT	ANOVA	—	<0.0001	ns	ns	ns	ns	ns	—
	ANCOVA	Soil moist.	<0.0001	ns	ns	ns	ns	ns	0.02
83 DAT	ANOVA	—	<0.0001	ns	ns	ns	ns	ns	—
	ANCOVA	Soil moist.	<0.0001	ns	ns	ns	ns	ns	ns
N ₂ O									
Rep. meas.	ANOVA	—	ns	<0.0001	<0.0001	ns	ns	<0.0001	—
48 DAT	ANOVA	—	ns	<0.0001	<0.0001	ns	ns	0.07	—
	ANCOVA	Soil moist.	ns	ns	<0.0001	ns	ns	0.07	0.04
69 DAT	ANOVA	—	ns	0.03	<0.0001	ns	ns	0.006	—
	ANCOVA	Soil moist.	ns	0.09	<0.0001	ns	ns	0.006	0.09
83 DAT	ANOVA	—	ns	0.09	<0.0001	ns	ns	0.06	—
	ANCOVA	Soil moist.	ns	ns	<0.0001	ns	ns	0.06	ns
		NO ₃ ⁻	ns	ns	0.01	ns	ns	0.06	0.03

^a Date effects in repeated measures ANOVA were always significant ($P<0.0001$). Block effects are not reported but were sometimes significant ($P<0.05$)

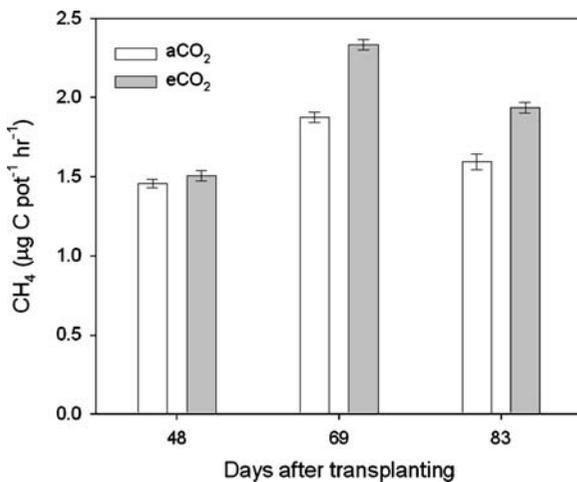


Fig. 3 Methane consumption on 48, 69, and 83 days after planting affected by CO₂ (aCO₂: ambient CO₂, eCO₂: elevated CO₂), averaged across water treatment and species composition treatment

emission was significantly related to soil moisture content, although the variability explained by soil moisture was small ($P < 0.0001$, $R^2 = 0.15$ Fig. 4b). The N₂O emission was not significantly related to pot respiration, and using pot respiration as a covariate in the ANCOVA did not change any of the treatment effects (data not shown). The N₂O emission 83 days after transplanting was not related to plant biomass, or plant biomass N, but was significantly related to NO₃⁻ concentrations in the soil ($P < 0.0001$, data not shown). The main water effect disappeared, and the species composition effect was less significant after adjusting for NO₃⁻ concentrations in the soil (Table 1).

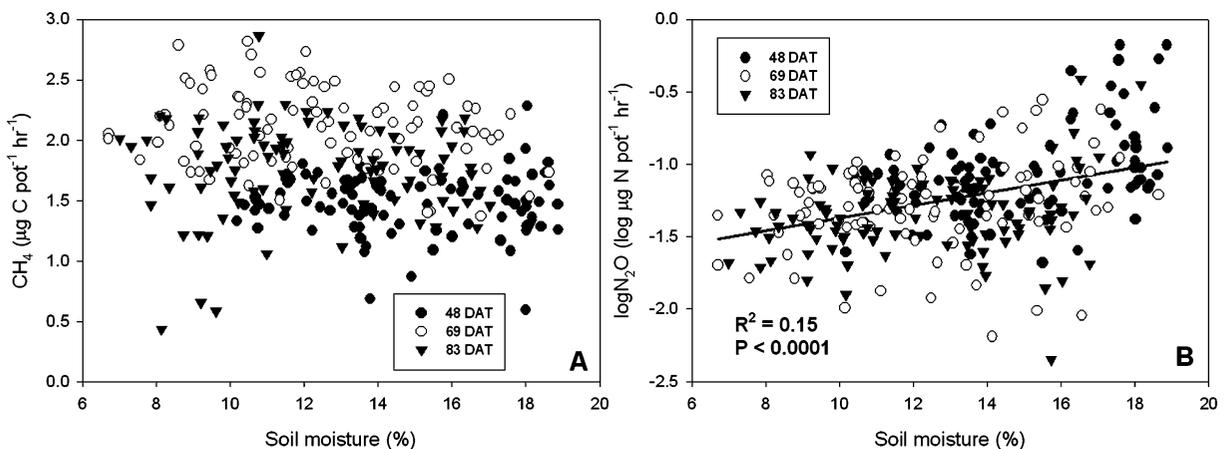


Fig. 4 **a** Methane consumption and **b** N₂O emission 48, 69, and 83 days after transplanting as a function of soil moisture

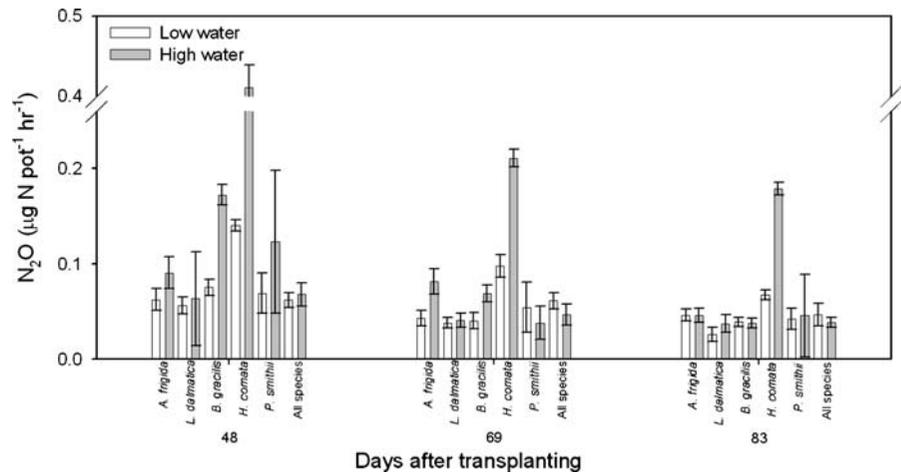
We observed no significant species number (1 vs. 5 species) effects on CH₄ consumption ($P = 0.12$, repeated measures ANOVA, data not shown). We did observe a significant reduction in N₂O emission with 5 species compared to 1 species ($P = 0.04$, repeated measures ANOVA), particularly under high water (species number*water interaction, $P = 0.05$). However, the species number effect and species number*water interaction disappeared during the last measurement, when plants were largest.

Discussion

Methane consumption significantly increased with elevated CO₂, but was not affected by soil water content or species composition, while N₂O emission was not affected by elevated CO₂, but significantly increased with increased water content and showed a significant species composition effect. The CO₂ effect on CH₄ consumption and the species composition effect on N₂O emission could not be explained by CO₂ or species composition effects on soil moisture.

We utilized a new method to measure microbially mediated CH₄ consumption and N₂O emission by eliminating effects of gas diffusivity. For CH₄, gas diffusivity may impede the influx of CH₄ into the soil when soil water content reaches a certain threshold (Del Grosso et al. 2000; Koschorreck and Conrad 1993). Mosier et al. (2008) reported that this threshold was around 10–20% water-filled pore space, or around 6.2–12% gravimetric soil moisture content (using a soil bulk density of 1 g cm⁻³ and a solid soil density of

Fig. 5 Nitrous oxide emission on 48, 69, and 83 days after transplanting affected by soil water content (low water: 15% soil moisture, high water: 20% soil moisture) and species composition averaged across the CO₂ treatment



2.65 g cm⁻³), for the USDA-ARS Central Plains Experimental Range where we collected our soil for this experiment. This threshold is at the lower end of soil moisture contents we maintained in our experiment. Thus gas diffusivity most likely limited CH₄ consumption in our pots during much of the time we did not circulate air through the pots, although diffusivity may have been higher in our pots than in the field due to gas exchange between the pot wall and the soil. Although we may not have completely eliminated gas diffusivity, we are confident that, during our measurements when we circulated air through the pots, we increased this threshold so that gas diffusivity was not the limiting factor for CH₄ consumption (and N₂O emission) anymore. If gas diffusivity played a significant role in CH₄ consumption during our measurements, we would have expected lower CH₄ consumption in the wetter soils under elevated CO₂, but we observed the opposite. We also did not observe a negative relationship between CH₄ consumption and soil moisture.

The activity of methanotrophs to oxidize CH₄ is limited by soil moisture content (Del Grosso et al. 2000; Stein and Hettiaratchi 2001). The increase in soil moisture under elevated CO₂ may have contributed to greater CH₄ consumption in pots under elevated CO₂ compared to pots with the same water treatment under ambient CO₂ (Fig. 2). However, the microbially mediated CH₄ consumption was not solely affected by soil moisture content, since CH₄ consumption under elevated CO₂ remained significantly higher after correcting for soil moisture. We also observed no relationship between CH₄ consump-

tion and soil moisture. These results suggest that the sensitivity of methanotroph activity to water content is not straightforward and that factors other than soil moisture contributed to the greater methanotroph activity under elevated CO₂.

High concentrations of NH₄⁺ and NO₃⁻ in the soil often inhibit methanotroph activity (Bédard and Knowles 1989), although increased CH₄ oxidation with NH₄⁺ fertilization in rice paddies has also been observed (Bodelier et al. 2000). Elevated CO₂ did not affect soil NO₃⁻ concentration (ANOVA, *P*=0.95) but increased the NH₄⁺ concentration by 25% (ANOVA, *P*=0.02; on average NH₄⁺ concentrations were 0.32 and 0.40 mg N kg⁻¹ soil for ambient and elevated CO₂ respectively). Although the increase in NH₄⁺ concentration under elevated CO₂ could potentially have reduced N limitation of methanotroph activity (Bodelier et al. 2000), the NH₄⁺ (or NO₃⁻) concentration in the soil was not related to CH₄ consumption.

It is possible that increased labile substrates under elevated CO₂ may have stimulated methanotrophs to consume CH₄. It is well known that elevated CO₂ can increase the input of labile C substrates into the soil (Cheng and Johnson 1998; Dijkstra et al. 2005; Hungate et al. 1997b). Evidence that methanotrophs are limited by C substrates is less clear. Conrad (1984) suggested that methanotrophs are able to use organic matter as a carbon source, and indeed several studies have suggested increased CH₄ consumption with greater availability of labile C substrates (Benstead et al. 1998; Goldman et al. 1995; Jacinthe and Lal 2005). However, elevated CO₂ reduced CH₄ consumption in a loblolly pine plantation that could not be explained by

potential increases in labile C substrates (Phillips et al. 2001b). We did not measure availability of labile C substrates in our experiment, but found a significant positive relationship between CH₄ consumption and respiration. Nevertheless, respiration could not explain the CO₂ treatment effect on CH₄ consumption, possibly because our respiration measurements may not have been a strong indicator of labile C substrates into the soil. A potentially greater availability of labile C substrates (as well as a greater soil moisture content) under elevated CO₂ may also have increased methanogen activity (i.e., production of CH₄). Although methanogen activity can occur in aerobic conditions (Khalil and Baggs 2005), soil moisture content in our experiment was most likely too small to cause a significant effect on CH₄ production (Von Fischer and Hedin 2007), and increased methanogen activity would have countered the CO₂ treatment effect on net CH₄ consumption.

In field studies a decrease in CH₄ consumption in response to elevated CO₂ has been associated with greater soil moisture content (Ambus and Robertson 1999; Ineson et al. 1998; McLain et al. 2002; McLain and Ahmann 2008). Rather than affecting methanotroph activity, it was suggested in these studies that soil moisture reduced CH₄ diffusivity into the soil thereby limiting CH₄ consumption or that soil moisture increased CH₄ production by methanogens. However, by eliminating the effects of CH₄ diffusivity on CH₄ consumption we were able to observe a significant increase in methanotroph activity under elevated CO₂, possibly due to an increase in soil moisture, and perhaps due to other mechanisms that are unclear. Our results indicate that elevated CO₂ effects on microbially mediated CH₄ consumption may be particularly important for ecosystems that are dry during much of the year and where CH₄ consumption is not generally constrained by gas diffusivity (Del Grosso et al. 2000). Indeed, CH₄ consumption in a semi-arid grassland in Colorado tended to be greater under elevated CO₂ than under ambient CO₂, despite an increase in soil moisture (Mosier et al. 2002). Field measurements from a free air CO₂ enrichment experiment in a semi-arid grassland in Wyoming showed greater CH₄ consumption under elevated CO₂ at times when soils were dry (i.e., during much of the growing season, Dijkstra et al., unpublished results).

Not surprisingly, soil moisture explained much of the higher N₂O emission in the high water treatment

(i.e., the water treatment effect disappeared or was strongly reduced when soil moisture was used as a covariate in the ANCOVA, Table 1). On the other hand, pots under elevated CO₂ had greater soil moisture contents during most of the experiment, but elevated CO₂ did not increase N₂O emission. Possibly, because of our frequent watering, the CO₂-induced soil moisture increase may not have been large enough to cause significant increases in N₂O emission that were observed in other studies (Arnone and Bohlen 1998; Kanerva et al. 2007; Kettunen et al. 2006; Robinson and Conroy 1999). It was suggested that a greater plant and microbial N demand under elevated CO₂ prevented an increase in N₂O emission (Hungate et al. 1997a; Mosier et al. 2002), or even reduced N₂O emission (Kettunen et al. 2007; Pleijel et al. 1998). However, total plant biomass N content at the end of the experiment was not significantly affected by the CO₂ treatment ($P > 0.1$, ANOVA, data not shown), while we do not have information about microbial N immobilization. Elevated CO₂ could also increase N₂O emission because of increased labile C input into the soil fueling the denitrifying community (Baggs et al. 2003a, b; Ineson et al. 1998; Kammann et al. 2008). However, nitrifiers that are an important contributor of N₂O emission in semi-arid systems (Mosier et al. 2008), do not require labile C as an energy source. Nitrification may have dominated the production of N₂O in our study, and thus N₂O emission may not have increased under elevated CO₂ despite potentially greater inputs of labile C substrates. Others have suggested that the lack of response of N₂O emission to elevated CO₂ was due to limited potential for denitrification (Ambus and Robertson 1999; Billings et al. 2002). The large variability in N₂O emission in response to elevated CO₂ among studies indicates a complex number of factors and interactions involved, which requires further investigation.

It is unclear what caused the large effect of species composition on N₂O emission. These effects could not be explained simply by species composition effects on soil moisture. When soil NO₃⁻ concentration was used as a covariate then species composition effects on N₂O emission were smaller, but remained significant (Table 1). The N₂O emission was largest under *H. comata*. This species also had the highest NO₃⁻ concentration in the soil at the end of the experiment (1.10 mg N kg⁻¹ compared to 0.67, 0.43,

0.74, 0.77, and 0.62 mg N kg⁻¹ for *A. frigida*, *L. dalmatica*, *B. gracilis*, *P. smithii*, and all species combined respectively). These results suggest that besides variation in soil moisture, variation in NO₃⁻ concentrations in the soil may have contributed to the variation in N₂O emission. However, other factors, such as root exudates and other root-derived available soil C (Baggs et al. 2003a, b; Ineson et al. 1998), differences in root morphology affecting root-microbial interactions (Brown et al. 1997), or differences in the composition of the nitrifying and denitrifying microbial community (Patra et al. 2006; Priha et al. 1999) may also have caused species composition effects on N₂O loss.

Increasing the plant species number from 1 to 5 did not affect CH₄ consumption and only reduced N₂O emission early in the experiment. Virtually nothing is known about plant species number or plant diversity effects on CH₄ fluxes and very little on N₂O fluxes. Reduced N₂O emission with increased plant species or plant functional diversity has been associated with more efficient plant N uptake (Epstein et al. 1998; Niklaus et al. 2006). More efficient N uptake early on in our experiment may also have contributed to the reduced N₂O emission in the mixed pots, which included species of diverse root morphologies.

Conclusion

Semi-arid grasslands have been shown to be very responsive to elevated CO₂ with large shifts in species composition and changes in C and N cycling (Dijkstra et al. 2008; Morgan et al. 2007; Pendall et al. 2003). Changes in species composition and in C and N cycling in response to elevated CO₂ likely also affect exchange of the greenhouse gases CH₄ and N₂O. However, the mechanisms involved in the microbially-mediated exchanges of CH₄ and N₂O between soils and the surrounding atmosphere have been obscured in previous experiments due to confounding effects of soil water content on microbial activity and diffusivity of gas through the soil matrix. We used a unique method that reduces soil matrix gas diffusional resistance to evaluate how elevated CO₂, soil water content, and plant species composition directly affect microbially mediated consumption of CH₄ and emission of N₂O. Elevated CO₂ significantly increased CH₄ consumption, while N₂O emission

significantly differed among plant species and significantly increased with greater soil moisture content. Although soil moisture is a key factor controlling biological processes in semi-arid grasslands, treatment effects on CH₄ consumption and N₂O emission often remained significant after correcting for treatment effects on soil moisture. This suggests that treatment-induced changes in soil and microbial properties such as availability of plant-derived labile C substrates, inorganic N, and microbial community composition are also important. More research is needed to better understand the interactions of environmental factors and different microbial groups contributing to CH₄ and N₂O exchange between the soil and atmosphere.

Acknowledgements We thank Joseph Hansen and Mary Smith for technical assistance. We thank Jean McLain, Joe von Fischer, and two anonymous reviewers for a critical review of a previous version of the manuscript. This publication is based upon work supported by the Agricultural Research Service under the ARS GRACENet Project.

References

- Ambus P, Robertson GP (1999) Fluxes of CH₄ and N₂O in aspen stands grown under ambient and twice-ambient CO₂. *Plant Soil* 209:1–8
- Ambus P, Zechmeister-Boltenstern S, Butterbach-Bahl K (2006) Sources of nitrous oxide emitted from European forest soils. *Biogeosciences* 3:135–145
- Arnone JA III, Bohlen PJ (1998) Stimulated N₂O flux from intact grassland monoliths after two growing seasons under elevated atmospheric CO₂. *Oecologia* 116:331–335
- Baggs EM, Blum H (2004) CH₄ oxidation and emissions of CH₄ and N₂O from *Lolium perenne* swards under elevated atmospheric CO₂. *Soil Biol Biochem* 36:713–723
- Baggs EM, Richter M, Cadisch G, Hartwig UA (2003a) Denitrification in grass swards is increased under elevated atmospheric CO₂. *Soil Biol Biochem* 35:729–732
- Baggs EM, Richter M, Hartwig UA, Cadisch G (2003b) Nitrous oxide emissions from grass swards during the eighth year of elevated atmospheric pCO₂ (Swiss FACE). *Glob Change Biol* 9:1214–1222
- Bédard C, Knowles R (1989) Physiology, biochemistry, and specific inhibitors of CH₄, NH₄⁺, and CO oxidation by methanotrophs and nitrifiers. *Microbiol Rev* 53:68–84
- Benstead J, King GM, Williams HG (1998) Methanol promotes atmospheric methane oxidation by methanotrophic cultures in soils. *Appl Environ Microbiol* 64:1091–1098
- Billings SA, Schaeffer SM, Evans RD (2002) Trace N gas losses and N mineralization in Mojave desert soils exposed to elevated CO₂. *Soil Biol Biochem* 34:1777–1784
- Bodelier PLE, Hahn AP, Arth IR, Frenzel P (2000) Effects of ammonium-based fertilisation on microbial processes

- involved in methane emission from soils planted with rice. *Biogeochemistry* 3:225–257
- Brown TN, Kulasiri D, Gaunt RE (1997) A root-morphology based simulation for plant/soil microbial ecosystem modeling. *Ecol Modell* 99:275–287
- Cheng W (1996) Measurement of rhizosphere respiration and organic matter decomposition using natural ^{13}C . *Plant Soil* 183:263–268
- Cheng W, Johnson DW (1998) Elevated CO_2 , rhizosphere processes, and soil organic matter decomposition. *Plant Soil* 202:167–174
- Conrad R (1984) Capacity of aerobic microorganisms to utilize and grow on atmospheric trace gases H_2 , CO_2 , CH_4 . In: Klug MJ, Reddy CA (eds) *Current perspectives in microbial ecology*. American Society for Microbiology, Washington, D.C. pp 461–467
- Crenshaw CL, Lauber C, Sinsabaugh RL (2008) Fungal control of nitrous oxide production in semiarid grassland. *Biogeochemistry* 87:17–27
- Del Grosso SJ, Parton WJ, Mosier AR, Ojima DS, Potter CS, Borken W, Brumme R, Butterbach-Bahl K, Crill PM, Dobbie K, Smith KA (2000) General CH_4 oxidation model and comparisons of CH_4 oxidation in natural and managed systems. *Glob Biogeochem Cycles* 14:999–1019
- Dijkstra FA, Hobbie SE, Reich PB, Knops JMH (2005) Divergent effects of elevated CO_2 , N fertilization, and plant diversity on soil C and N dynamics in a grassland field experiment. *Plant Soil* 272:41–52
- Dijkstra FA, Pendall E, Mosier AR, King JY, Milchunas DG, Morgan JA (2008) Long-term enhancement of N availability and plant growth under elevated CO_2 in a semi-arid grassland. *Funct Ecol* 22:975–982
- Epstein HE, Burke IC, Mosier AR, Hutchinson GL (1998) Plant functional type effects on trace gas fluxes in the shortgrass steppe. *Biogeochemistry* 42:145–168
- Frank DA, Groffman PM (1998) Denitrification in a semi-arid grazing ecosystem. *Oecologia* 117:564–569
- Galbally IE, Kirstine WV, Meyer CP, Wang YP (2008) Soil-atmosphere trace gas exchange in semiarid and arid zones. *J Environ Qual* 37:599–607
- Goldman MB, Groffman PM, Pouyat RV, McDonnell MJ, Pickett STA (1995) CH_4 uptake and N availability in forest soils along an urban to rural gradient. *Soil Biol Biochem* 27:281–286
- Goverde M, Erhardt A (2003) Effects of elevated CO_2 on development and larval food-plant preference in the butterfly *Coenonympha pamphilus* (Lepidoptera, Satyridae). *Glob Change Biol* 9:74–83
- Heijmans MMPD, Klees H, De Visser W, Berendse F (2002) Response of a *Sphagnum* bog plant community to elevated CO_2 and N supply. *Plant Ecol* 162:123–134
- Hungate BA, Lund CP, Pearson HL, Chapin FS III (1997a) Elevated CO_2 and nutrient addition alter soil N cycling and N trace gas fluxes with early season wet-up in a California annual grassland. *Biogeochemistry* 37:89–109
- Hungate BA, Holland EA, Jackson RB, Chapin FS III, Mooney HA, Field CB (1997b) The fate of carbon in grassland under carbon dioxide enrichment. *Nature* 388:576–579
- Hütsch BW, Webster CP, Powlson DS (1994) Methane oxidation in soil as affected by land use, soil pH and N fertilization. *Soil Biol Biochem* 26:1613–1622
- Huxman TE, Smith MD, Fay PA, Knapp AK, Shaw MR, Lolk ME, Smith SD, Tissue DT, Zak JC, Weltzin JF, Pockman WT, Sala OE, Haddad BM, Harte J, Koch GW, Schwinnig S, Small EE, Williams DG (2004) Convergence across biomes to a common rain-use efficiency. *Nature* 429:651–654
- Ineson P, Coward PA, Hartwig UA (1998) Soil gas fluxes of N_2O , CH_4 and CO_2 beneath *Lolium perenne* under elevated CO_2 : The Swiss free air carbon dioxide enrichment experiment. *Plant Soil* 198:89–95
- Intergovernmental Panel on Climate Change (2007) Working Group I Report. The Physical Science Basis. Technical Summary. Available online. <http://www.ipcc.ch>. Accessed 7 May 2009
- Jacinthe PA, Lal R (2005) Labile carbon and methane uptake as affected by tillage intensity in a Mollisol. *Soil Till Res* 80:35–45
- Kammann C, Müller C, Grünhage L, Jäger H-J (2008) Elevated CO_2 stimulates N_2O emissions in permanent grassland. *Soil Biol Biochem* 40:2194–2205
- Kanerva T, Regina K, Rämö K, Ojanperä K, Manninen S (2007) Fluxes of N_2O , CH_4 and CO_2 in a meadow ecosystem exposed to elevated ozone and carbon dioxide for three years. *Environ Poll* 145:818–828
- Kang H, Freeman C, Ashendon TW (2001) Effects of elevated CO_2 on fen peat biogeochemistry. *Sci Total Environ* 279:45–50
- Kettunen R, Saarnio S, Martikainen PJ, Silvola J (2005) Elevated CO_2 concentration and nitrogen fertilisation effects on N_2O and CH_4 fluxes and biomass production of *Phleum pratense* on farmed peat soil (2005) *Soil Biol Biochem* 37:739–750
- Kettunen R, Saarnio S, Martikainen PJ, Silvola J (2006) Increase of N_2O fluxes in agricultural peat and sandy soil under elevated CO_2 concentration: concomitant changes in soil moisture, groundwater table and biomass production of *Phleum pratense*. *Nutr Cycl Agroecosyst* 74:175–189
- Kettunen R, Saarnio S, Martikainen PJ, Silvola J (2007) Can a mixed stand of N_2 -fixing and non-fixing plants restrict N_2O emissions with increasing CO_2 concentration? *Soil Biol Biochem* 39:2538–2546
- Khalil MI, Baggs EM (2005) CH_4 oxidation and N_2O emission at varied soil water-filled pore spaces and headspace CH_4 concentrations. *Soil Biol Biochem* 37:1785–1794
- Koschorreck M, Conrad R (1993) Oxidation of atmospheric methane in soil: measurements in the field, in soil cores and in soil samples. *Glob Biogeochem Cycles* 7:109–122
- McLain JET, Ahmann D (2008) Increased moisture and methanogenesis contribute to reduced methane oxidation in elevated CO_2 soils. *Biol Fert Soils* 44:623–631
- McLain JET, Kepler TB, Ahmann DM (2002) Belowground factors mediating changes in methane consumption in a forest soil under elevated CO_2 . *Glob Biogeochem Cycles* 16:1050. doi:10.1029/2001GB001439
- Menyailo OV, Huwe B (1999) Activity of denitrification and dynamics of N_2O release in soils under six tree species and grassland in central Siberia. *J Plant Nutr Soil Sci* 162:533–538
- Menyailo OV, Hungate BA (2003) Interactive effects of tree species and soil moisture on methane consumption. *Soil Biol Biochem* 35:625–628

- Morgan JA, Mosier AR, Milchunas DG, LeCain DR, Nelson JA, Parton WJ (2004a) CO₂ enhances productivity, alters species composition, and reduces digestibility of shortgrass steppe vegetation. *Ecol Appl* 14:208–219
- Morgan JA, Pataki DE, Körner C, Clark H, Del Grosso SJ, Grunzweig JM, Knapp AK, Mosier AR, Newton PCD, Niklaus PA, Nippert JB, Nowak RS, Parton WJ, Polley HW, Shaw MR (2004b) Water relations in grassland and desert ecosystems exposed to elevated atmospheric CO₂. *Oecologia* 140:11–25
- Morgan JA, Milchunas DG, LeCain DR, West M, Mosier AR (2007) Carbon dioxide enrichment alters plant community structure and accelerates shrub growth in the shortgrass steppe. *P Natl Acad Sci USA* 104:14724–14729
- Mosier A, Schimel D, Valentine D, Bronson K, Parton W (1991) Methane and nitrous oxide fluxes in native, fertilized and cultivated grasslands. *Nature* 350:330–332
- Mosier AR, Parton WJ, Valentine DW, Ojima DS, Schimel DS, Heinemeyer O (1997) CH₄ and N₂O fluxes in the Colorado shortgrass steppe 2. Long-term impact of land use change. *Glob Biogeochem Cycles* 11:29–42
- Mosier AR, Morgan JA, King JY, LeCain D, Milchunas DG (2002) Soil-atmosphere exchange of CH₄, CO₂, NO_x, and N₂O in the Colorado shortgrass steppe under elevated CO₂. *Plant Soil* 240:201–211
- Mosier AR, Parton WJ, Martin RE, Valentine DW, Ojima DS, Schimel DS, Burke IC, Adair EC, Del Grosso SJ (2008) Soil-atmosphere exchange of trace gases in the Colorado shortgrass steppe. In: Lauenroth WK, Burke IC (eds) *Ecology of the shortgrass steppe*. Oxford University Press, Oxford, pp 342–372
- Niklaus PA, Wardle DA, Tate KR (2006) Effects of plant species diversity and composition on nitrogen cycling and the trace gas balance of soils. *Plant Soil* 282:83–98
- Norton U, Mosier AR, Morgan JA, Derner JD, Ingram LJ, Stahl PD (2008) Moisture pulses, trace gas emissions and soil C and N in cheatgrass and native grass-dominated sagebrush-steppe in Wyoming, USA. *Soil Biol Biochem* 40:1421–1431
- Patra AK, Abbadie L, Clays-Josserand A, Degrange V, Grayston SJ, Guillaumeaud N, Loiseau P, Louault F, Mahmood S, Nazaret S, Philippot L, Poly F, Prosser JI, Roux XL (2006) Effects of management regime and plant species on the enzyme activity and genetic structure of N-fixing, denitrifying and nitrifying bacterial communities in grassland soils. *Environ Microbiol* 8:1005–1016
- Pendall E, Del Grosso S, King JY, LeCain DR, Milchunas DG, Morgan JA, Mosier AR, Ojima DS, Parton WJ, Tans PP, White JWC (2003) Elevated atmospheric CO₂ effects and soil water feedbacks on soil respiration components in a Colorado grassland. *Global Biogeochem Cy* 17: doi:10.1029/2001GB001821
- Phillips RL, Whalen SC, Schlesinger WH (2001a) Influence of atmospheric CO₂ enrichment on methane consumption in a temperature forest soil. *Glob Change Biol* 7:557–563
- Phillips RL, Whalen SC, Schlesinger WH (2001b) Response of soil methanotrophic activity to carbon dioxide enrichment in a North Carolina coniferous forest. *Soil Biol Biochem* 33:793–800
- Pleijel H, Sild J, Danielsson H, Klemetsson L (1998) Nitrous oxide emissions from a wheat field in response to elevated carbon dioxide concentration and open-top chamber enclosure. *Environ Pollut* 102, S1:167–171
- Potts DL, Huxman TE, Cable JM, English NB, Ignace DD, Eilts JA, Mason MJ, Weltzin JF, Williams DG (2006) Antecedent moisture and seasonal precipitation influence the response of canopy-scale carbon and water exchange to rainfall pulses in a semi-arid grassland. *New Phytol* 170:849–860
- Priha O, Grayston SJ, Pennanen T, Smolander A (1999) Microbial activities related to C and N cycling and microbial community structure in the rhizospheres of *Pinus sylvestris*, *Picea abies* and *Betula pendula* seedlings in an organic and mineral soil. *FEMS Microbiol Ecol* 30:187–199
- Robinson D, Conroy JP (1999) A possible plant-mediated feedback between elevated CO₂, denitrification and the enhanced greenhouse effect. *Soil Biol Biochem* 31:43–53
- Stein VB, Hettiaratchi JPA (2001) Methane oxidation in three Alberta soils: influence of soil parameters and methane flux rates. *Environ Technol* 22:101–111
- Ullah S, Frasier R, King L, Picotte-Anderson N, Moore TR (2008) Potential fluxes of N₂O and CH₄ from soils of three forest types in Eastern Canada. *Soil Biol Biochem* 40:986–994
- Von Fischer JC, Hedin LO (2007) Controls on soil methane fluxes: tests of biophysical mechanisms using stable isotope tracers. *Glob Biogeochem Cycles* 21, doi:10.1029/2006GB002687